Electrostatics in Wind-Blown Sand

Jasper F. Kok^{1,2,*} and Nilton O. Renno^{1,2}

¹Applied Physics Program, University of Michigan, Ann Arbor, Michigan 48109, USA

²Atmospheric, Oceanic, and Space Sciences, University of Michigan, Ann Arbor, Michigan 48109, USA

(Received 5 July 2007; published 8 January 2008)

Wind-blown sand, or "saltation," is an important geological process, and the primary source of atmospheric mineral dust aerosols. Significant discrepancies exist between classical saltation theory and measurements. We show here that these discrepancies can be resolved by the inclusion of sand electrification in a physically based saltation model. Indeed, we find that electric forces enhance the concentration of saltating particles and cause them to travel closer to the surface, in agreement with measurements. Our results thus indicate that sand electrification plays an important role in saltation.

DOI: 10.1103/PhysRevLett.100.014501

Introduction.—In wind-blown sand, or "saltation," sand grains are propelled by the wind while they bounce along the surface (Fig. 1). Saltation plays a key role in various geological processes, including wind erosion, sediment transport, and the formation of sand dunes [1,2]. Moreover, the impact of saltating grains on the soil surface is the primary emission mechanism of mineral dust aerosols [2]. These small airborne dust particles affect the Earth system in several ways, including by scattering and absorbing radiation [3], acting as cloud seeds [4], and providing nutrients to ecosystems [5].

Pioneering measurements by Schmidt *et al.* [6] show that electric fields (E fields) of up to \sim 160 kV/m can be generated in saltation under moderate wind conditions. Laboratory experiments show that such E fields facilitate the lifting of sand particles by winds, and can even directly lift sand particles from the surface [7]. Large E fields are also predicted to occur in Martian saltation, possibly producing large quantities of hydrogen peroxide and making the Martian surface inhospitable to life as we know it [8].

It is known from industrial handling of powders that particle collisions tend to leave smaller particles with net negative charge and larger particles with net positive charge [9]. The physical mechanism governing this charge transfer is poorly understood, although various studies suggest that asymmetric rubbing (i.e., a small area of the small particle rubs over a large area of the large particle) causes a net transfer of electrons from the larger to the smaller particle [10]. We hypothesize that a similar process occurs in saltation, where particles bounce along a soil surface that can be interpreted as the surface of an infinitely large particle (Fig. 1). Saltating particles are thus expected to charge negatively with respect to the surface, as is indeed indicated by measurements of upward-pointing E fields in wind tunnel and field experiments [6,11]. The occurrence of upward-pointing E fields in dust devils and dust storms [12] also suggests negatively charged particles over a positively charged surface.

In this Letter, we present the first physically based numerical model of saltation that includes the generation of electric fields and the effects of electric forces on PACS numbers: 47.55.Kf, 41.20.Cv, 45.70.Mg, 92.40.Gc

saltation. We show that recent measurements in saltation [13–15] cannot be explained by classical saltation theory [1,16], but are consistent with the predictions of our model when sand electrification is included.

Theory.—Saltation is initiated when the wind shear stress τ exceeds the threshold value τ_t necessary to move surface particles. The height-integrated particle mass flux Q is a good measure of the intensity of saltation. Experiments show that Q increases approximately cubically with wind shear velocity, $u* = \sqrt{\tau/\rho_a}$ [1,2]; that is,

$$Q = Q_0(\rho_a/g)u^{*3},$$
 (1)

where ρ_a is the air density, g is the gravitational acceleration, and Q_0 is the dimensionless particle mass flux.

Model description.—A detailed description of our model is given in Ref. [17], but we describe it here briefly. We model saltation as the interplay of four processes [18]: (i) the motion of saltating particles, (ii) the modification of the wind profile through momentum transfer with saltating particles, (iii) the collisions of particles with the soil sur-

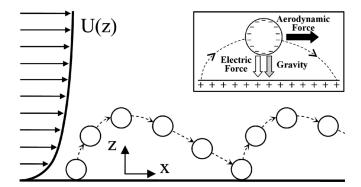


FIG. 1. Schematic of saltation, showing the logarithmic wind profile U(z) to the left of an idealized spherical sand particle propelled by the wind and bouncing along the surface. After lift-off from the surface, saltating particles obtain horizontal momentum from the wind, which is partially converted into vertical momentum after colliding with the surface and rebounding. The inset shows the force diagram of a negatively charged saltating particle over the positively charged soil surface.

face, and (iv) the lifting of surface grains by wind stress and by particles impacting the soil surface. The main innovation of our model over previous models (e.g., [18]) is that the charging of saltating particles during collisions with other particles and the surface is included. The effect of the resulting electric forces on particle motion and the threshold shear stress (τ_t) is explicitly accounted for [7,17].

We model particle motion in two dimensions by considering gravitational, fluid, and electric forces (Fig. 1). The effects of turbulence and midair collisions on particle trajectories are neglected because these effects are relatively small for typical shear velocities [2,18]. Results of laboratory and numerical studies are used to model the collision of particles with the soil surface, including the ejection of surface grains [18].

For the calculation of the wind profile, we make the classical assumption that, in steady-state saltation, the fluid shear stress at the surface stays at the threshold value (τ_t) necessary to initiate the motion of surface particles [16,17]. The particle concentration per unit area is then given by [17]

$$N = \tau_p(0) / \overline{\tau_{\rm sp}(0)} = (\tau - \tau_t) / \overline{\tau_{\rm sp}(0)},$$
 (2)

where $\tau_p(0)$ is the total shear stress exerted by saltating particles at the surface, and $\overline{\tau_{\rm sp}(0)}$ is the average surface stress exerted by a single saltating particle, as computed from the particle trajectories. The size distribution of saltating particles is assumed to be similar to that of the parent soil for particles of 100–500 μ m, in agreement with field measurements [17].

Sand electrification.—The model accounts for electrostatic charging of saltating particles during collisions with other particles [17] and the surface. Although collisional charge transfer between grains of granular material is observed in a variety of physical systems [6,9,11,12], the charging mechanism is not well understood. Nonetheless, Desch and Cuzzi [19] proposed a model in which the collisional charge transfer depends on the preexisting charges, the particle sizes, and the difference in the particles' contact potential. They proposed that

$$q_S' = C_1(q_S + q_L) - C_2 \Delta \Phi,$$

$$q_L' = (1 - C_1)(q_S + q_L) + C_2 \Delta \Phi,$$
(3)

where q_S and q_L are the charges of the smaller and larger particles before the collision, q_S' and q_L' are the charges after the collision, $\Delta\Phi$ is the difference in particle contact potential, and C_1 and C_2 are functions of the mutual capacitances (and thus the radii) of the two particles, as defined by Eqs. 5–10 of Ref. [19]. For particles of similar composition (i.e., $\Delta\Phi=0$), such as typical soil particles, (3) suggests that no charge transfer occurs when the colliding particles are not initially charged, which contradicts observations [6,9,11,12]. To mitigate this problem, we propose an effective contact potential difference between

particle pairs of similar composition but different sizes. That is,

$$\Delta\Phi_{\rm eff} = S(r_L - r_S)/(r_L + r_S),\tag{4}$$

where S (in volts) is a physical parameter that scales the collisional charge transfer, and r_S and r_L are the radii of the small and large particles. This simple model has a functional form consistent with observations—smaller particles acquire net negative charge during collisions with larger particles, and the charge transfer is reduced as the relative difference in particle size decreases. Since saltating particles impacting the soil surface tend to interact with multiple surface grains [1,18], we interpret the soil surface as the surface of an infinitely large particle (i.e., $r_L = \infty$). By calibrating the model with E-field measurements in saltation [6], we found that $S = 6 \pm 4$ V (see Fig. 2).

The soil surface is assumed to be conducting, both because charge exchange with saltating particles provides charge mobility, and because conducting films of water are generally adsorbed on soils [7]. Since the height to which particles saltate is generally much smaller than the horizontal extent over which saltation occurs, we use the infinite plane approximation to determine the electric field E from the calculated space charge density ρ_c and soil surface charge density σ [7],

$$E(z) = \frac{1}{2\varepsilon_0} \left(\sigma + \int_0^z \rho_c(z') dz' - \int_z^\infty \rho_c(z') dz' \right), \quad (5)$$

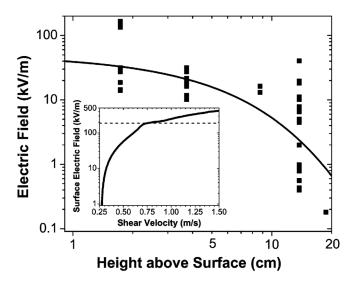


FIG. 2. Comparison of measured (squares) and modeled (solid line) E fields in saltation as a function of height. The measurements [6] were taken for winds of 4–12 m/s at 1.5 m height, which was estimated to correspond to an average shear velocity of 0.5 ± 0.1 m/s [2]. The soil particle size distribution was taken as typical for the broad top of a dune [22], where the measurements were made. The inset shows the dependence of the surface E field on shear velocity for the size distribution reported in Ref. [13], with the dashed line corresponding to the electric lifting threshold [7].

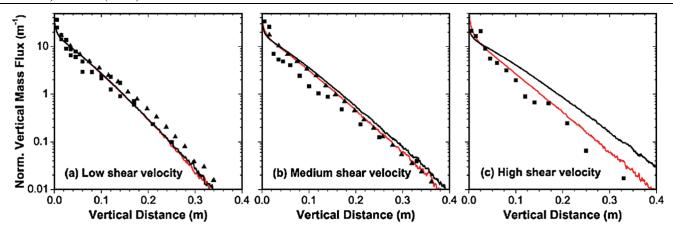


FIG. 3 (color). Vertical profiles of saltation mass flux measured in field experiments (squares [13] and triangles [14]), and compared to model results with and without electric forces (solid red lines and solid black lines, respectively). Both measured and modeled mass flux profiles are normalized by their total mass flux to simplify comparison. Results are shown for (a) low shear velocity ($u^* = 0.32 \text{ m/s}$), (b) medium shear velocity ($u^* = 0.47 \text{ m/s}$), and (c) high shear velocity ($u^* = 0.63 \text{ m/s}$). Model results were obtained for the size distribution reported in Ref. [13].

where ε_0 is the electric permittivity of air. The effect of the surface electric field on the threshold shear velocity (u_t^*) is calculated using Eq. 9 in Ref. [7].

Model results.—Comparisons of vertical and horizontal profiles of the mass flux in saltation show that predictions of our model are in agreement with measurements (Fig. 3 and Ref. [17]). To the best of our knowledge, the model presented here is the first physically based model capable of accurately reproducing observed particle mass flux profiles. For high shear velocity, the agreement with field measurements improves when electric forces are included [Fig. 3(c)]. The height-integrated mass flux predicted by our model is also in good qualitative agreement with wind tunnel results [17].

Our model predicts that E fields increase sharply towards the surface and with wind speed (Fig. 2), in agreement with measurements [6,11]. The surface E field is of particular interest because E fields larger than $\sim 80 \text{ kV/m}$ reduce significantly the wind shear stress necessary to lift surface particles [7]. Our model finds that this effect contributes to an approximate doubling of the particle concentration due to sand electrification at large shear velocities (Fig. 4).

In addition to increasing the concentration of saltating particles, electric forces also affect particle trajectories. A characteristic height of saltation is the height z_{50} below which 50% of the mass transport occurs. Classical saltation theory [1,16] predicts that increases in wind speed produce increases in the momentum of saltating particles, causing them to impact and rebound from the surface at higher speed, and therefore reach larger heights. However, recent measurements show that z_{50} stays approximately constant as the wind speed increases [13–15]. This clear discrepancy between measurements and theory can be resolved by the inclusion of sand electrification in our physically based saltation model (Fig. 5). As the negatively charged saltat-

ing particles bounce along the positively charged surface (Fig. 1) [6,11], the downward electric force causes particles to travel closer to the surface and at reduced horizontal speed [20]. Since the downward electric force increases with wind speed (Fig. 2), z_{50} remains approximately constant up to moderate shear velocities, in good agreement with measurements [13–15]. At larger shear velocities, electric forces become strong enough to lower the threshold shear velocity [7,17], which reduces the near-surface winds and thus z_{50} . We plan to test this prediction with future field measurements.

Conclusions.—We developed the first physically based numerical saltation model that includes the effects of sand

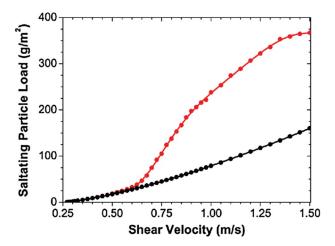


FIG. 4 (color). Mass load of saltating particles modeled with and without electric forces (red circles and black circles, respectively) as a function of shear velocity, for the size distribution reported in Ref. [13]. Electric forces cause the saltation particle load to increase at a given shear velocity by reducing the wind stress required to lift surface particles [7] and by reducing the average surface stress exerted by a single saltating particle [20].

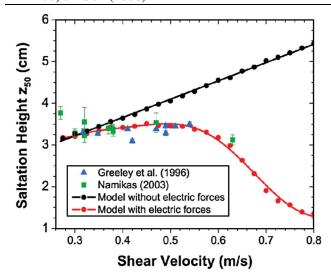


FIG. 5 (color). Dependence of the characteristic saltation height z_{50} (see text) on the wind shear velocity. Classical saltation theory predicts that z_{50} increases strongly with shear velocity [1,16], which our model also predicts when electric forces are not included (black circles). However, field measurements (squares [13] and triangles [14]) show that z_{50} remains approximately constant. Inclusion of sand electrification in our model (red circles) resolves this discrepancy. The linear increase in z_{50} without sand electrification is consistent with results from an independent numerical model [23]. Values of z_{50} were obtained from Refs. [13,14] by fitting an exponential function to the measured vertical mass flux profiles, as described in Ref. [13]. Error bars represent the uncertainty in the fitting parameters. Model results were obtained for the size distribution reported in Ref. [13].

electrification. Significant discrepancies exist between classical saltation theory and field measurements [13]. We show that the inclusion of sand electrification in saltation models can resolve these discrepancies.

Model results show that sand electrification increases the particle concentration at a given wind shear velocity [7,20] (Fig. 4). Moreover, the downward electric force on saltating particles lowers their trajectories, improving the agreement between model predictions and measurements [13–15] (Fig. 5). Our results thus indicate that sand electrification plays an important role in saltation.

We are also investigating the effect of the classical assumption that the shear stress at the surface remains at the threshold value for particle entrainment [16,17]. Initial results suggest that removing this assumption could also explain some of the discrepancies between theory and measurements [21].

Although the current study focuses on terrestrial saltation, our model predicts that electric discharges occur during intense Martian saltation [24].

We thank Scott Schmidt for providing his original *E*-field measurements, and Michael Bretz, William Kuhn, Shanna Shaked, Earle Williams, Daniel Lacks, and an anonymous reviewer for comments. Finally, we thank NSF for financial support through Grant No. ATM 0622539.

*jfkok@umich.edu

- [1] R.A. Bagnold, *The Physics of Blown Sand and Desert Dunes* (Methuen, London, 1941).
- [2] Y. P. Shao, *Physics and Modelling of Wind Erosion* (Kluwer Academic, Dordrecht, 2000).
- [3] I. Tegen and A. A. Lacis, J. Geophys. Res. 101, 19237 (1996).
- [4] P. J. DeMott et al., Geophys. Res. Lett. 30, 1732 (2003).
- [5] T. D. Jickells *et al.*, Science **308**, 67 (2005).
- [6] D. S. Schmidt, R. A. Schmidt, and J. D. Dent, J. Geophys. Res. 103, 8997 (1998); D. S. Schmidt (private communication).
- [7] J. F. Kok and N. O. Renno, Geophys. Res. Lett. 33, L19S10 (2006).
- [8] S. K. Atreya et al., Astrobiology 6, 439 (2006).
- [9] I. I. Inculet, G. S. P. Castle, and G. Aartsen, Chem. Eng. Sci. 61, 2249 (2006).
- J. Lowell and W. S. Truscott, J. Phys. D 19, 1281 (1986);
 D. J. Lacks and A. Levandovsky, J. Electrost. 65, 107 (2007).
- [11] X. J. Zheng, N. Huang, and Y. H. Zhou, J. Geophys. Res. 108, 4322 (2003); J. J. Qu et al., Sci. China Ser. D 47, 529 (2004).
- [12] G. D. Freier, J. Geophys. Res. 65, 3504 (1960); C. D. Stow, Weather 24, 134 (1969).
- [13] S. L. Namikas, Sedimentology **50**, 303 (2003).
- [14] R. Greeley, D.G. Blumberg, and S.H. Williams, Sedimentology **43**, 41 (1996).
- [15] K. R. Rasmussen and M. Sorensen, J. Geophys. Res. (to be published).
- [16] P. R. Owen, J. Fluid Mech. 20, 225 (1964).
- [17] See EPAPS Document No. E-PRLTAO-99-032753 for a detailed model description, as well as model results of the horizontal mass flux profile and the height-integrated mass flux. For more information on EPAPS, see http://www.aip.org/pubservs/epaps.html.
- [18] R.S. Anderson and P.K. Haff, Acta Mech. Suppl. 1, 21 (1991).
- [19] S. J. Desch and J. N. Cuzzi, Icarus 143, 87 (2000).
- [20] The decrease in particle speed due to electric forces reduces the average surface stress exerted by a single saltating particle $[\tau_{sp}(0)]$ in (2), which contributes to the increase in particle concentration seen in Fig. 4.
- [21] J. F. Kok and N. O. Renno (to be published).
- [22] N. Lancaster, J. Sediment. Petrol. 56, 395 (1986).
- [23] M. P. Almeida, J. S. Andrade, and H. J. Herrmann, Phys. Rev. Lett. **96**, 018001 (2006).
- [24] N. O. Renno and J. F. Kok (to be published).

Vol 451|14 February 2008

NEWS & VIEWS



GRANULAR MATTER

Static in motion

Troy Shinbrot and Hans J. Herrmann

Wind-blown desert sands can charge up spontaneously. But although sand flow and the forces on charged bodies are well studied separately, surprisingly little is known of what happens when the two combine.

The strange things that happen when granular media and electrostatic charge mix have long exercised the minds of both scientists and engineers. In the latest episode in this saga, Kok and Renno, writing in *Physical Review Letters*¹, focus on how static charges affect the aeolian — wind-borne — transport of sand. As the basic ingredient of sand encroachment and dune motion (Fig. 1), this is a process of more than academic interest, with implications for desertification and coastal management.

But first some history. As long ago as 1867, William Thomson, better known as Lord Kelvin, studied the charging of both falling grains and water droplets in his attempts to understand the origin of atmospheric lightning². Eighty years on from that, physicist E. W. B. Gill drew on his observations of sparking and radio interference on the Macedonian front during the First World War to produce a similar effect in the laboratory³.

Grains in an electric field are also known to self-assemble into complex patterns⁴. Other poorly understood, but bothersome, phenomena can be traced back to particle charging, too. Clouds of charged dust regularly produce

devastating explosions in grain and coal plants⁵. On the Moon, charged grit attaches itself to spacesuits and works its way into suit joints, causing them to leak air and so cut exploration time. The problem was such that Apollo 16 commander John Young considered dust the number one concern in returning there. That sentiment was echoed in NASA's most recent study on Mars exploration⁶.

Back on Earth, aeolian flows have been measured in both field⁷ and laboratory⁸ experiments. These investigations showed that charges acquired by wind-blown sand can be sufficient to levitate the grains, and even to violently eject them (Fig. 2, overleaf). Despite experimental data demonstrating that grains readily acquire substantial charges, with few notable exceptions⁸, surprisingly little fundamental analysis has been done on how charge affects granular flow, and several essential questions remain to be answered.

First, it has been established that even carefully prepared, identical materials charge one another ^{9,10}. But how exactly does this work? In a desert environment, where wind-blown sand grains have little but other, similar grains to

rub against, yet still manage to acquire charge, this is a significant puzzle. Kok and Renno¹ go some way to finding an answer, constructing an effective charging relation, for pairs of particles of the same composition but different sizes, that describes the empirical fact that smaller sand grains tend to charge negatively, and larger grains positively.

Even so, the mechanism for the underlying charge transfer remains at best tenuously understood. One proposal is that smaller particles rise to the top of a sand cloud, where they encounter more highly mobile, negative ions in the air, and so themselves become negatively charged¹. Another theory holds that asymmetric particle collisions heat the smaller particles more effectively than the larger ones, possibly leading to a transfer of charge between them¹⁰.

A second unanswered question is how the charge that is measured on sand grains affects aeolian flow. Kok and Renno produce a model for the motion of wind-blown sand ('saltation') that takes electrostatic interactions into account by assuming charged sand grains to be attracted to Earth, which acts as an infinitely large, oppositely charged particle. This causes

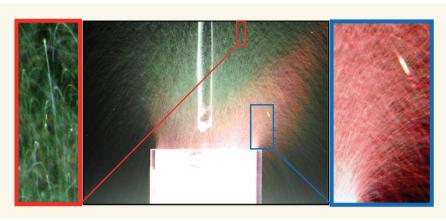


Figure 2 | Granular fountain. In this laboratory demonstration (centre) of the explosive potential of charged grains, glass beads 500 micrometres in diameter are charged by repeatedly pouring them through a vertical acrylic tube into an acrylic container. The charges on the grains become so large that the beads cannot remain at rest, and they spontaneously form a fountain that erupts from the container, even after the inflow has ceased. If the tube is electrically grounded, these ejections do not occur. Left inset, trajectories of individual grains; right inset, an apparent example of a granular aggregate coexisting alongside individual grains. Kok and Renno¹ produced a model of the development of such charge on sand. (Colours are digitally enhanced; image and experiment courtesy of F. K. Wittel, ETH-Zürich.)

charged sand grains to fall back down faster than they would do otherwise. At the same time, the authors note that charged grains tend to levitate (or even jump considerable distances, as seen in Figure 2), and so should become more easily airborne than neutral grains. Both of these apparently contradictory propositions seem sound, but their combined effect on sand flow, and the effect of charges of both signs on a saltating bed, remain to be fully understood. Further progress will require

a concerted combination of careful measurements and simulations rarely attempted to date⁸.

One final, extremely basic, question is: will charged sand grains attract or repel each other? Both phenomena are observed. If grains simply acquire a uniform charge, they will of course repel each other according to Coulomb's law. But dipole moments — a result of an uneven distribution of electric charge — on single particles have also been observed, and local

assemblies of charge are known to form at points of contact between grains¹¹. Consequently, similarly treated grains are seen to both repel each other and to adhere in clusters³ in the same system (Fig. 2). How a collection of grains in, say, a desert sandstorm will behave will ultimately depend on the charge distributions on individual grains, which are poorly understood.

It is remarkable that, since Kelvin's time, so basic a question as how a common material such as sand becomes charged and flows remains unresolved. Work such as that of Kok and Renno¹ demonstrates how complex the research challenges are. The simple combination of grains and charge will no doubt be generating more surprises in the years to come.

Troy Shinbrot and Hans J. Herrmann are at the Institut für Baustoffe, ETH-Zürich, Schafmattstrasse 6, CH-8093 Zürich, Switzerland. e-mail: hans@ifb.baug.ethz.ch

- 1. Kok, J. F. & Renno, N. O. Phys. Rev. Lett. **100,** 014501 (2008).
- Thomson, W. Proc. R. Soc. Lond. 16, 67-72 (1867).
- 3. Gill, E. W. B. Nature 162, 568-569 (1948).
- Sapozhnikov, M. V., Tolmachev, Y. V., Aranson, I. S. & Kwok, W. K. Phys. Rev. Lett. 90, 114301 (2003).
- Palmer, K. N. Dust Explosions and Fires (Chapman & Hall, London, 1973).
- Calvin, W. et al. Rep. 2013 Mars Science Orbiter (MSO) Second Science Analysis Group http://mepag.jpl.nasa.gov/ reports/index.html (2007)
- Schmidt, D. S., Schmidt, R. A. & Dent, J. D. J. Geophys. Res. 103, 8897–9001 (1998).
- Zheng, X. J., Huang, N. & Zhou, Y. H. J. Geophys. Res. 108, 4322 (2003).
- 9. Shaw, P. E. Proc. Phys. Soc. 39, 449-452 (1927).
- 10. Latham, J. Q. J. R. Meteorol. Soc. 90, 91-95 (1964).
- Terris, B. D., Stern, J. E., Rugar, D. & Mamin, H. J. Phys. Rev. Lett. 63, 2669–2672 (1989).